

## Using the ARP-03 for high-resolution mapping of calcic horizons\*\*

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**Abstract.** The aim of this work is to present a fast and cheap method for high-resolution mapping of calcic horizons in vineyards based on geoelectrical proximal sensing. The study area, 45 ha located in southern Sicily (Italy), was characterized by an old, partially dismantled marine terrace and soils with a calcic horizon at different depths. The geoelectrical investigation consisted of a survey of the soil electrical resistivity recorded with the Automatic Resistivity Profiling-03 sensor. The electrical resistivity values at three pseudo-depths, 0-50, 0-100 and 0-170 cm, were spatialized by means of ordinary kriging. A principal component analysis of the three electrical resistivity maps was carried out. During the survey, 18 boreholes, located at different electrical resistivity values, were made for soil description and sampling. The depth to the calcic horizon showed a strong correlation with electrical resistivity. The regression model between calcic horizon and the principal component analysis factors with the highest correlation coefficients was selected to spatialise the calcic horizon values. An Normalized Difference Vegetation Index map was used to validate the calcic horizon map in terms of crop response to different soil rooting depths. The strengths of this method are the quick, non-invasive kind of survey, the relevance for vine vigour, and the high spatial resolution of the final map.

**Key words:** precision viticulture, soil conservation, irrigation, Mediterranean, geophysics

### INTRODUCTION

Understanding functional soil features at high spatial resolution is of utmost importance for 'site-specific' crop management. In Mediterranean countries, the depth of the upper boundary of the calcic horizon (Dk), described as Bck or Ck, usually corresponds to the soil rooting depth, and thus plays a key role in determining soil fertility and crop water availability. These horizons are usually brought near the surface by erosion processes or incorrect land levelling and

land preparation for the vineyard, which can lead to severe soil degradation (Martínez-Casasnovas and Ramos, 2006) and even desertification (Sánchez *et al.*, 2004).

For this reason, a high-detail map of soils and soil depth is necessary for correct land preparation, vineyard husbandry and soil conservation practices within vineyard (Uson *et al.*, 1998). In addition, this kind of investigation is essential for proper and economically sustainable planning of irrigation, tailored to the spatial variability of soil water retention capacity and soil depth.

In particular, it is well known that both the quantity and quality of grapes in the Mediterranean region are significantly determined by water supply (Costantini *et al.*, 2006; Deloire *et al.*, 2004). Plant water stress is primarily conditioned by rainfall and crop management, but also by rooting depth and soil water retention capacity (Costantini *et al.*, 2010). Variations in soil properties, often associated with a variation in topography, appear to be a crucial factor in variability of vine vigour (Bramley and Lamb, 2003).

Bramley and Lanyon (2002) demonstrated that yield variation in Australian vineyards with a 'Terra Rossa' soil (red clayey soil) over limestone was driven by disparity in the water available in the root zone, which in turn was controlled by topography variations. Moreover, Bramley and Hamilton (2004) showed that the pattern of the yield variation was closely correlated with soil properties, much more than the influence of seasonal differences.

In the last decade, on-the-go measurement of the spatial variability of the apparent soil electrical conductivity (ECa), or resistivity (ER) has become an invaluable tool to map physical and hydrological parameters of soils. Originally used to investigate soil salinity (Doolittle *et al.*, 2001), this technology has also been used in the last two decades in new investigative fields related to physical and hydrological parameters of soils.

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In recent years, proximal sensing methods have also been used in preliminary studies for precision viticulture. Taylor *et al.* (2009) used proximal sensing methods and spatial statistics to replace traditional soil maps of vineyards with new high-resolution soil maps. Bourennane *et al.* (2012) used the principal component analysis (PCA) of electrical resistivity (ER) data to estimate the soil water content in a vineyard.

The aim of the present study was to validate a fast and cheap method for high-resolution soil mapping based on proximal sensing and factorial analysis. The method combined the results provided by a geophysical proximal sensor with those from the description of a limited number of borings to obtain a high spatial resolution map of the calcic horizon depth in the studied vineyards. The strengths of the method are the quickness and the limited manpower needed to obtain a very high-resolution soil map of the vineyards.

#### MATERIALS AND METHODS

The study area was inside one of Sicily largest wineries, 'Villa Albius', situated between Acate and Marina d'Acate (Ragusa) in south-eastern Sicily (Fig. 1). The farm covers about 1 000 hectares, 800 of which are used for irrigated viticulture. It produces indigenous varieties such as Nero d'Avola, Inzolia, Grillo, Frappato, as well as international varieties such as Chardonnay, Syrah and Cabernet Sauvignon. For this study, ten vineyards (in total 45 ha) were chosen for a detailed proximal sensing survey. The main cultivar in these vineyards was Nero d'Avola, although two small vineyards of Frappato were also present.

Agriculture is the main business in the area and involves frequent soil cultivation to control weeds to prevent water and nutrient competition. The vineyards are located on a broad Plio-Pleistocene marine terrace, a widespread geomorphologic unit in this area of Sicily, as in many other Mediterranean environments. This terrace is characterized by a first level of calcarenite, a few meters thick and dating to the

Early-Middle Pleistocene, overlying yellow sands partially cemented by calcium carbonate attributed to the Early Pleistocene (Grasso, 2000). The terrace is deeply cut by fluvial valleys and strongly eroded, especially along the outer edges.

Some of the vineyards investigated in this study are situated on flat surface of the terrace, characterized by reddish soils developed on the calcarenite. The vineyards situated in the eastern part of the study area, on the slopes of the terrace edge, exhibited either strongly eroded yellowish soils (loose and poorly weathered) developed on the yellow sands, or thin reddish soils developed on the patches of calcarenite residues.

The slopes are very irregular, being cut by small valleys with NW-SE and SE-NW directions, and have different aspects. Moreover, the slopes usually show strong erosion, sometimes in the form of rills and gullies. In contrast, in the vineyards on the terrace, the relatively flat morphology limited soil erosion, with the exception of the small rises, where moderate sheet erosion was active.

The 1:250 000 soil map of Sicily (Fantappiè *et al.*, 2010) reported the Soil Typologic Units (STUs) of the marine terraces of the study area: soil region 62.2, described as hills of Sicily on clayey flysch, limestone, sandstone and gypsum, and coastal plains with Mediterranean subtropical climate (Costantini *et al.*, 2004). In the studied vineyards, the soil map of Sicily reported Endoleptic Calcisols (STU 62.2CLha1) on the slopes of the terrace and Haplic Cambisols (STU 62.2CMeu1) on the flat surface of the terrace.

Boundary conditions of the experiment (climate, topography, vineyard management, rootstock, *etc.*) could be considered representative of the study area and typical of Mediterranean viticulture on soils formed from homogeneous bedrock of marine calcareous rocks.

The soil proximal survey of the vineyards was carried out with the ARP-03 (Automatic Resistivity Profiling), a mobile soil electrical resistivity mapping system consisting of a multiprobe system of three arrays, provided with a GPS and pulled by an ATV vehicle. The ARP-03, produced since

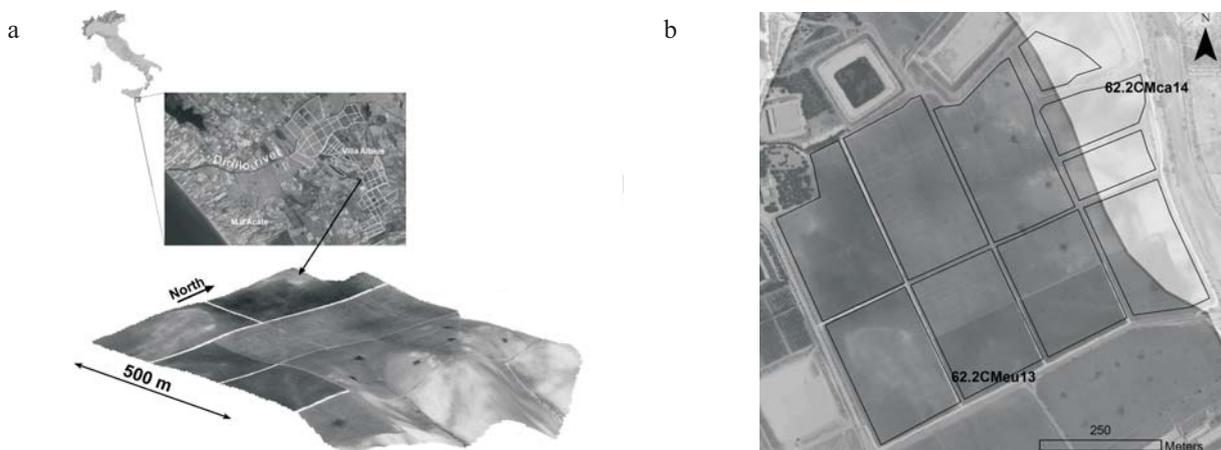


Fig. 1. Study area – a, extract of the soil map of Sicily – b (1:250 000).

2001, represents the innovation of the MuCEP (Multi Continuous Electrical Profiling) prototype (Panissod *et al.*, 1997; Tabbagh *et al.*, 2000). Both instruments were conceived by Geocarta (France), a spin-off society of the CNRS (National Scientific Research Centre, France), for applications in archaeology (Papadopoulos *et al.*, 2009) and precision agriculture (Michot *et al.*, 2007).

The ARP-03 device consisted of pairs of toothed metal wheels operating as injection electrodes and three pairs of toothed metal wheels functioning as receivers and measuring the difference in electrical potential. The shape of the toothed wheels assured contact between the electrodes and the soil matrix. The distance between each pair of receivers was conceived and calibrated to investigate three soil pseudo-depths, about 0-50 (ER<sub>50</sub>), 0-100 (ER<sub>100</sub>) and 0-170 cm (ER<sub>170</sub>). The measurements of apparent electrical resistivity (ER) in  $\Omega\text{m}$  and the altitude ( $z$ ) were georeferenced with a GPS and recorded on a field laptop computer.

During the survey, the ARP-03 swaths were spaced about 20 m apart, every 8-9 vine rows, and the total ER data numbered about 4 500 for each pseudo-depth *ie* about 100 measurements per hectare. This was a good compromise between spatial accuracy and costs, as demonstrated by Andrenelli *et al.* (2010) and Costantini *et al.* (2009). The rapidity of the survey was one of the strengths of this methodology: the proximal survey and the borings for soil description and sampling were conducted by only two people in a single day, 23 March 2010.

The ER data recorded with the ARP-03 were interpolated by ordinary kriging with a Gaussian variogram model, using the ArcGIS 9.3 geostatistical analysis (ESRI, Inc., California, USA). The lag distance was 10 m and the maximum range 120 m. In the same way, the altitude data ( $z$ ), measured by GPS, were used to prepare a detailed 1 x 1 m grid Digital Elevation Model (DEM) of the vineyards.

On the day of the ARP-03 survey, 18 soil borings were carried out with a hand auger up to 1.2 m depth. The location of the borings was chosen on the basis of the ARP-03 results to provide reasonable coverage of the ER values. The soils at the boring sites were described according to NRCS-USDA-NSSC (1998) and 13 of them were sampled for the laboratory analysis. Particular attention was given to textural changes, pedofeatures (clay coatings, CaCO<sub>3</sub> nodules and concretions), and depth of the horizon boundaries. All soils showed a poorly structured calcic horizon at different depths, described as Bck or Ck. The boundary between the calcic horizon and the overlying Ap, Bw or Bt horizons was always clear. The Dk was very variable. In the shallowest and most eroded soils, a Ck horizon outcropped at the surface.

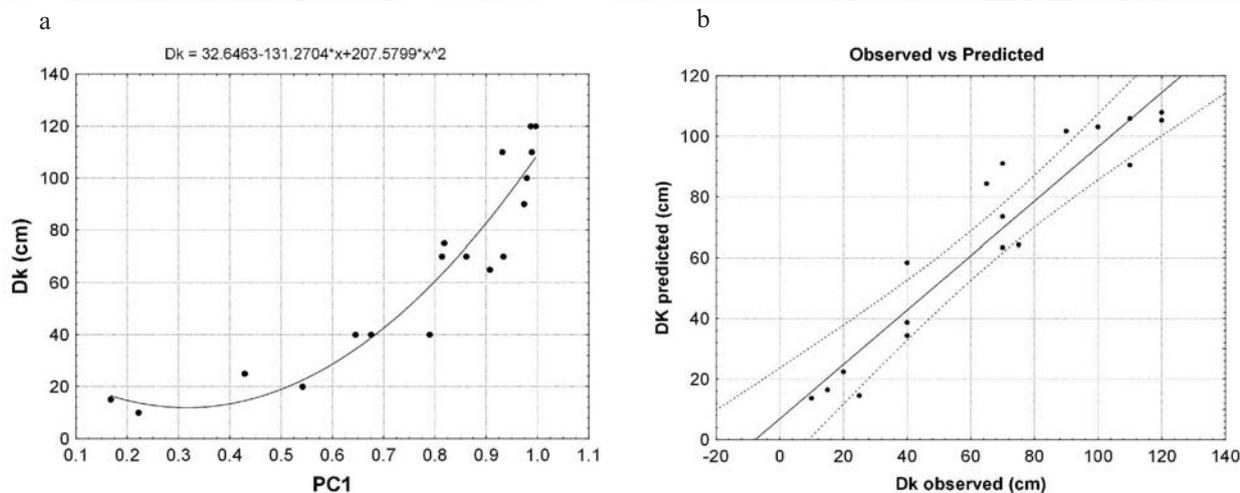
Laboratory analyses were performed on 10 samples (Ap and subsurface horizons) to characterize the soils according to the official Italian methods (Mi.P.A.F., 2000). Soil texture was determined with a sieve and hydrometer. CaCO<sub>3</sub> concentration was measured gas-volumetrically by

the addition of HCl in a Dietrich-Frühling calcimeter. EC was measured in 1:2 soil/water solution. Organic carbon was determined by the Walkley-Black procedure. Soil water content ( $\theta_w$ ) was determined by the thermo-gravimetric method after drying at 105° C for 24 h. The available water capacity (AWC) was estimated with SPAW software (<http://hydrolab.arsusda.gov>) which applies the prediction equations of Saxton and Rawls (2006). The laboratory analysis, together with the field descriptions, were used to classify the soil borings according to WRB (IUSS Working Group WRB, 2006) and to group them into Soil Typological Units (STUs).

To eliminate multicollinearity of the three ER maps (ER<sub>50</sub>, ER<sub>100</sub> and ER<sub>170</sub>) prior to the regression analysis, a Principal Component Analysis (PCA) was performed by SAGA-Gis, using the variance-covariance matrix. Bourenane *et al.* (2012) used this method in the same way for electrical resistivity data. PCA converts a set of possible correlated variables into a set of uncorrelated values called principal components, namely PCs (Jolliffe, 1986). The PCs are ordered by decreasing explained variance and most of the variance is usually explained by the first few PCs. In our case, we only needed to use the map of PC1, which explained 98.15% of the variance. The map of PC1 was standardized to the range of 0-1 by subtracting the minimum value and dividing by the maximum one in order to avoid negative values. The standardized PC1 values corresponding to the boring locations were extracted to perform the regression analysis with the measured Dk values. Linear, power, logarithmic, exponential and polynomial regression models were elaborated to identify the best fitting model to correlate Dk to PC1. A polynomial model was used for the regression analysis (Fig. 2). The regression model was used to interpolate the Dk on the basis of the PC1.

To validate the results of the Dk regression modelling, we used a Normalized Difference Vegetation Index (NDVI) map created with satellite remote sensing as the response variable. In fact, it is well known that there is a strong influence of soil depth on plant vigour. An NDVI map is a quick synthetic estimate of spatial variability of the Leaf Area Index (LAI) and thus of crop vigour (Johnson *et al.*, 2003). A GeoEye-1 multispectral satellite image was acquired to calculate the NDVI map of the vineyards. The image was acquired with the satellite sensor on June 23, 2010, during the veraison period of the grapes. The GeoEye-1 satellite sensor was developed by GeoEye (U.S.) and launched in September 2008. It is capable of acquiring image data at a resolution of 0.41 m and 1.65 m, respectively, for panchromatic (B and W) and multispectral images. The Imagine software package (ERDAS Inc., Atlanta, USA) was used for image processing. The NDVI real number outputs (from -1 to +1) were scaled and converted to a byte-binary format 0-255 by the software.

Lamb *et al.* (2004) demonstrated the utility of airborne and/or satellite images collected during the growing season for the prediction of characteristics and variations of grape-



**Fig. 2.** Polynomial model of: a – Dk spatial regression and b – scatterplot of Dk-observed vs. Dk-predicted with the linear correlation fit and 95% confidence.

vine phenology. According to Bramley and Lamb (2003), the range time of  $\pm 2$  weeks from veraison is the best predictor of fruit colour and phenolics at vintage. Since the veraison period for the Nero d'Avola cultivar is in the middle or in the second half of July, we acquired a satellite image at the end of June.

After spatialization of the Dk values and validation with the NDVI, a high-detail Soil Typological Units (STUs) map was created, grouping the Dk map into four classes. The Dk limits used for the STUs map were chosen on the basis of the soil typologies identified from the borings.

## RESULTS

The ARP-03 provided 4 448 soil electrical resistivity (ER) values for three depths (0-50, 0-100 and 0-170 cm). The interpolated maps of the ER data showed a root mean square error (RMSE) of 25.5, 16.9, 40.5  $\Omega\text{m}$  for ER<sub>50</sub>, ER<sub>100</sub> and ER<sub>170</sub> respectively. The mean values of the ER maps are reported in Table 1. The maps of the data interpolation are reported in Fig. 3, while the descriptive statistics of the ER data are in Table 1.

The differences in absolute values were clearly visible, especially in the high resistivity areas. ER<sub>100</sub> gave the lowest mean value (82.6  $\Omega\text{m}$ ) and the lowest standard deviation (93.9  $\Omega\text{m}$ ). The ER maps showed very similar patterns and were highly correlated among themselves, with  $R^2$  between 0.916 (ER<sub>50</sub> with ER<sub>100</sub>) and 0.935 (ER<sub>50</sub> with ER<sub>170</sub>). The PC1 explained 98.15% of the total variance (Table 2). The PC1 map is presented in Fig. 4.

The borehole soil observations revealed a poorly structured, sandy, calcic horizon at variable depths overlying the bedrock, namely the fractured limestone or the partially cemented yellow sands (R or Cr horizon). In some cases, the boreholes were not deep enough to detect the bedrock which was deeper than 1.2 m. In contrast, the calcic horizon (BCK or Ck) was identified in all the soil borings, either over the limestone or on the yellow sands. The BCK or Ck horizon was clearly recognisable from its white-yellowish colour, loose structure and frequent CaCO<sub>3</sub> nodules and soft concentrations. The upper boundary of the calcic horizon was always clear and characterized by sudden richness of calcareous nodules and concretions and by a strong reduction of soil structure. The differentiation between BCK and Ck was

**Table 1.** Descriptive statistics of altitude (h), ER<sub>50</sub>, ER<sub>100</sub> and ER<sub>170</sub>

Parameter	Mean	n	Min	Max	25 Percentile	75 Percentile	Standard deviation
h (m)	102.0	4448	86.6	107.7	101.4	103.8	3.3
ER <sub>50</sub> ( $\Omega\text{m}$ )	161.9	4448	18.2	985.1	59.4	209.8	159.3
ER <sub>100</sub> ( $\Omega\text{m}$ )	82.6	4448	4.6	652.1	22.1	106.3	93.9
ER <sub>170</sub> ( $\Omega\text{m}$ )	245.2	4448	15.9	1492.3	72.0	312.3	249.1

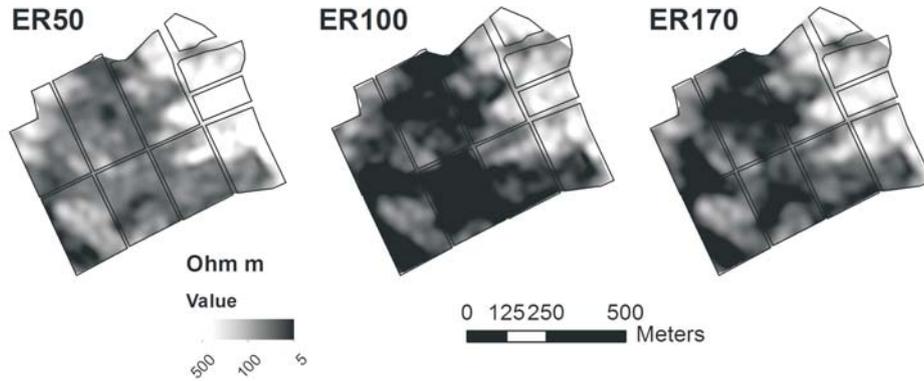


Fig. 3. Interpolated maps of ER<sub>50</sub>, ER<sub>100</sub> and ER<sub>170</sub>.

Table 2. Statistics of the three principal components

Parameter	PC1	PC2	PC3
ER <sub>50</sub>	0.506	-0.808	0.301
ER <sub>100</sub>	0.296	-0.165	-0.941
ER <sub>170</sub>	0.810	0.565	0.156
Eigenvalue	38634742067	572127498	156528955
% Variance	98.15	1.45	0.4
Cumulative variance	98.15	99.6	100

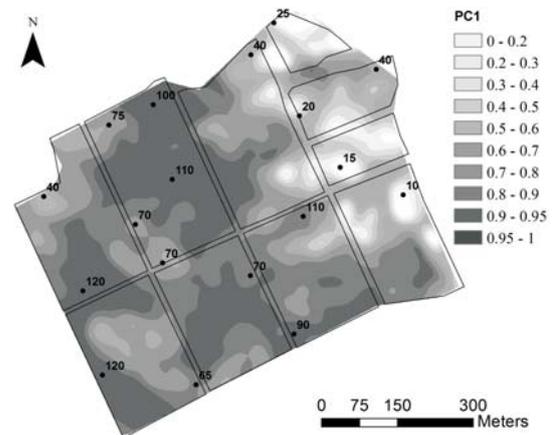


Fig. 4. PC1 map. The dots represent the location of borings with the Dk values (cm).

based on the soil morphological stage: poor (Bck) or absent (Ck) soil structure and presence/absence of the colour change in the parent material. The Bck horizon often passed to the Ck horizon with a gradual or diffuse boundary.

The Pearson correlation coefficients between the Dk and ER data were very high: 0.87, 0.93 and 0.90 ( $p < 0.01$ ) for ER<sub>50</sub>, ER<sub>100</sub> and ER<sub>170</sub>, respectively. The Pearson coefficient between Dk and PC1 was also very high: 0.92 ( $p < 0.01$ ).

The best regression model between Dk and PC1 (equation) was the polynomial (Fig. 3) which showed a determination coefficient  $R^2$  0.90 ( $p < 0.01$ ):

$$Dk = 32.6423 - (131.2704 PC1) + (207.5799 PC1^2)$$

The root mean square error (RMSE) of the model was 11.5 cm, in relation to a mean Dk value of 75 cm. The model was used to spatialise the Dk values within the studied vineyards, resulting in the Dk map in Fig. 5.

The cross-correlation between the NDVI and Dk values (Figs 5 and 6) was highly significant ( $R^2 = 0.89$ ,  $p < 0.001$ ). The NDVI map showed generally low values in the NE slope, characterized by the most eroded soils, and the highest values in the slight depressions on the flat surface of the terrace where deeper and less eroded soils were more frequent. The comparison between the NDVI and Dk maps allowed us to test the influence of soil depth on vine vigour.

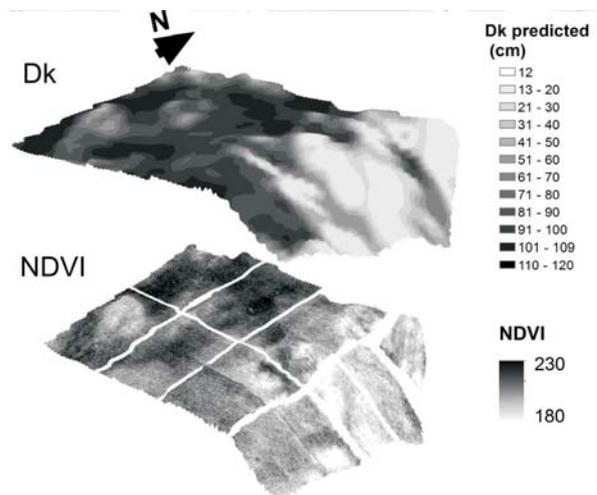


Fig. 5. Dk and NDVI maps overlying the DEM.

The soil observations made during the survey were grouped into Soil Typological Units (STUs) in order to create a high-detail soil map. On the basis of the field description and laboratory analysis, four STUs (Fig. 7) were identified and classified as follows (IUSS Working Group WRB, 2006):

- STU A: Cutanic Luvisol (Profondic, Chromic). This STU represented the less eroded soils, characterized by a calcic horizon deeper than 100 cm. The soil matrix of the Ap, Bw and Bt horizons was almost completely decarbonated ( $\text{CaCO}_3$ : 0.5%), although some  $\text{CaCO}_3$  hard nodules and some calcarenite fragments were present. The sandy loam texture in the superficial horizons (Ap and Bw) became sandy clay loam in the Bt horizon. The mean AWC was  $80 \pm 3 \text{ mm m}^{-1}$ .
- STU B: Haplic Calcisol (Arenic, Chromic). This STU showed a calcic horizon (BCK) deeper than 75 cm and the bedrock was about 100-120 cm deep. Also in this case the soil matrix of the Ap and Bw horizons was decarbonated ( $\text{CaCO}_3$ : 0.5-1%) with few calcarenite fragments and

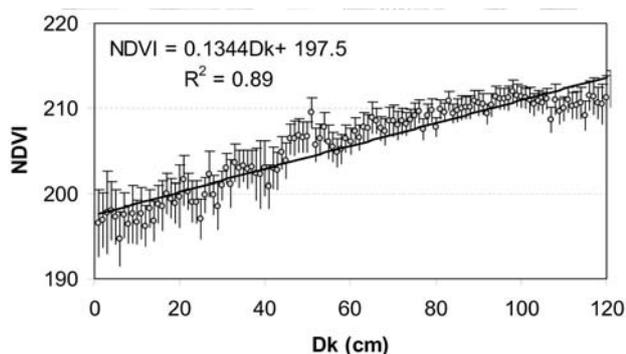


Fig. 6. Linear correlation between NDVI and Dk. The error bars show the -90% and +90% quantiles.

- $\text{CaCO}_3$  hard nodules. The texture ranged from loamy sand to sandy loam, and the mean AWC was  $57 \pm 4 \text{ mm m}^{-1}$ .
- STU C: Endoleptic Calcisol (Arenic). The upper limit of the calcic horizon in this STU was shallow (between 25 and 75 cm) and the fractured bedrock was found between 70 and 100 cm. The  $\text{CaCO}_3$  content in the Ap horizon was also high, from 10 to 20%, and  $\text{CaCO}_3$  hard nodules and calcarenite fragments were common. The  $\text{CaCO}_3$  soft concentrations and nodules became frequent in the BCK horizon, while  $\text{CaCO}_3$  was about 25%. The BCK horizon often changed gradually into the underlying Ck horizon, poorly structured and pale yellowish in colour. The texture was sandy, and the mean AWC was  $47 \pm 4 \text{ mm m}^{-1}$ .
- STU D: Aric Regosol (Calcaric, Arenic). This STU represented the most eroded soils in the vineyards. The calcic horizon was not distinct from the Ap and C horizons because all the horizons had been mixed by deep ploughing. In some cases, a shallow Ap horizon, with poor structure and pale brown in colour, was present, but the Bw soil horizon was absent. The fractured bedrock or the partially cemented unweathered sands were shallower than 50-70 cm.  $\text{CaCO}_3$  concentration of the Ap/Ck horizon was higher than 18% and  $\text{CaCO}_3$  soft concentrations and nodules were common. The texture was sandy and the mean AWC was  $39 \pm 4 \text{ mm m}^{-1}$ .

The pH ( $\text{H}_2\text{O}$ ) was sub-alkaline in all soil samples (7.8-8.5) and the EC varied from about 100 to  $188 \mu\text{S cm}^{-1}$  in the superficial horizons (Ap) and from about 30 to  $160 \mu\text{S cm}^{-1}$  in the deeper ones (Bw, BCK, *etc.*). These values in the superficial and deeper horizons excluded the presence of significant soil salinity. The organic carbon was very low in all the soils ( $< 10 \text{ g kg}^{-1}$ ). The soil water content ( $\theta$ ) was slightly lower than the field capacity for all the samples and ranged between 1.2 and 11%. As expected, the highest moisture content (8-11%) was always found in the deepest horizons of STUs A and B, while the lower ones were found

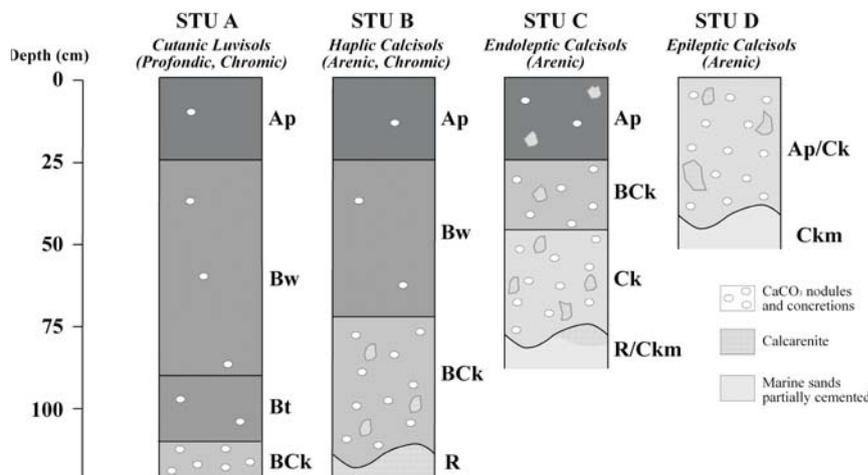
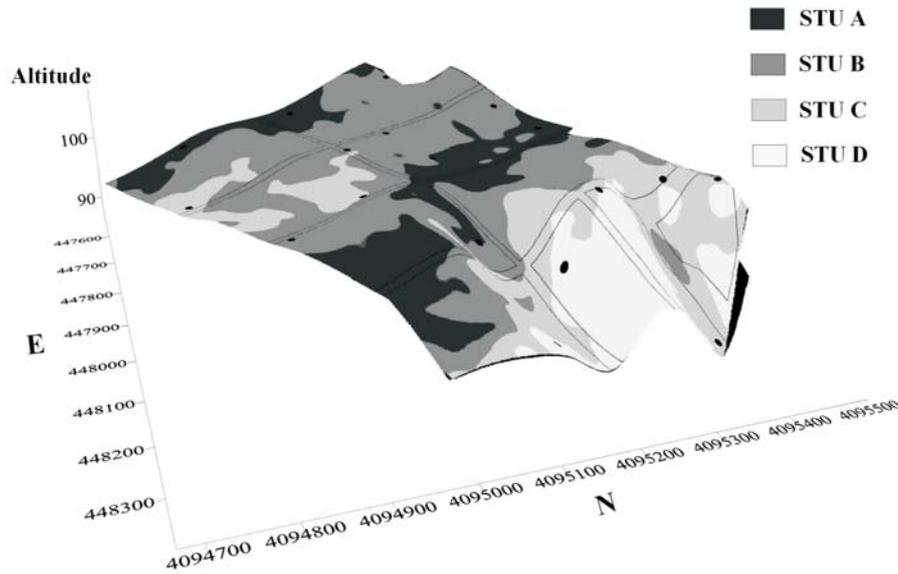


Fig. 7. Schematic description and classification of the STUs.



**Fig. 8.** STUs map obtained from the Dk prediction map and soil boring descriptions. The black dots are the boring sites.

in the samples of STU D. The variability of calcic horizon depth, strongly related to soil depth and conservation, was the soil feature that more characterized the STUs variability.

The STUs map was developed on the basis of the Dk prediction map (Fig. 8). The Dk limits between STUs was chosen on the basis of the modal value of the calcic horizon depth of each STU, measured during the soil boring descriptions. The depth of calcic horizons was below 100 cm for STU A, between 75 and 100 cm for STU B, between 25 and 75 for STU C and shallower than 25 cm for STU D.

DISCUSSION

The ER maps showed very similar patterns, although ER<sub>100</sub> had lower mean values and lower standard deviations. This was probably due to higher soil water content in the deeper soil horizons, usually located between 30-50 and 90-120 cm. The ER distribution within the study area was characterized by a more resistive area on the slope of the terrace (North-East of the map) and by less resistive areas on the terraced surface. Although having a similar pattern, the ER<sub>170</sub> cm map showed higher mean values (243.5 Ωm) than ER<sub>50</sub> (160 Ωm) and ER<sub>100</sub> (81.7 Ωm), indicating a high resistivity layer in depth. Moreover, the highest ER values were measured on the slopes, where the strongly eroded soils were located. In contrast, the lowest values were typically measured in the flat or slightly concave areas, in the dark reddish soils that were well-structured and with very scarce gravel.

The PCA confirmed the high correlation and the multicollinearity of the ER values. In fact, PC1 explained 98.15% of the variance, while PC2 and PC3 did not substantially improve the analysis.

The non-linear polynomial regression of Dk with PC1 can be clearly explained by the very strong difference between the electrical resistivity of the upper soil horizons (Ap and Bw) and the bedrock. When the bedrock and the calcic horizon were as shallow as the ARP-03 electrical signal penetration depth, the ER quickly became high. Therefore, a 50 cm deep soil showed an ER value 8-10 times higher than a soil with similar texture but 100 cm deep. This characteristic alone allowed to make a high-resolution spatialization of the Dk within the studied vineyards.

The comparison between the NDVI and Dk maps permitted to test the soil map in terms of the relationship between depth to the calcic horizon and vine vigour. The correlation between the NDVI and Dk values (Figs 5 and 6) was highly significant ( $R^2 = 0.89$ ,  $p < 0.01$ ). Although the NDVI map could not directly validate the Dk predicted map, the strong correlation between these maps supported the accuracy of the spatialization of the Dk values. In fact, the B<sub>ck</sub> and C<sub>k</sub> horizons represented the deepest part of the soil, just overlying the bedrock and with poor structure, low organic matter, low fertility and low water availability. The deeper the calcic horizon, the higher the soil fertility and water retention.

The field description of soil borings and laboratory analysis allowed us to recognise four STUs, distinguished mainly on the basis of the depth to the calcic horizon, strongly related to texture and CaCO<sub>3</sub>. The results of the other soil analyses, namely pH, EC and OC, were not statistically different between the four STUs. This is due to the great homogeneity of the area in terms of parent material and land use.

## CONCLUSIONS

1. The demonstrated the possibility to use ARP-03 proximal sensors and a few soil boring descriptions for a high-detail map of calcic horizon depth (Dk) in vineyards of SE Sicily (Italy). The Dk prediction map represented the spatial variability of the calcic horizon, closely related to soil depth and conservation.

2. The significant correlation between the Dk and NDVI maps demonstrated a strong relationship between vine vigour and soil depth in the study area.

3. A high-detail Soil Typological Units (STUs) map, derived from the Dk map and a few soil boring descriptions, can be created with a limited amount of fieldwork.

4. The resulting maps of the STUs as discrete variable and Dk as continuous variable can be used to plan a diversified field management in terms of fertilization, pruning and irrigation.

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